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UPPER MIocene PRE-TERRIGENOUS DEPOSITS OF THE QUEGLIA UNIT (ABRUZZO, CENTRAL APENNINES, ITALY): BIOSTRATIGRAPHIC DATA

Abstract - An integrated biostratigraphic study (calcareous nannofossils, planktic and benthic foraminifera) of the Upper Miocene pre-terrigenous deposits of the Queglia unit was carried out. The section studied (Pescosansonesco, Central Apennines-Abruzzo) is representative of a deepening-upward foreland ramp occupying a relatively external position with respect to the well-known Laga foredeep basin in the Messinian.

Amaurolithus primus and *Globorotalia gr. miotumida* were observed at the base of the section, where a gradual change from shallow to deeper water shelf was recognized. In the overlaying hemipelagic sediments, FO of *Bulimina echinata*, FO and FCO of *Turborotalia multiloba*, and FO of *Amaurolithus amplificus* were recorded.

These events enabled us to recognize of the planktic foraminifera *Turborotalia multiloba* Subzone and the calcareous nannofossil *Calcidiscus leptoporus* MA Subzone.

Key words - Biostratigraphy, Calcareous Nannofossils, Foraminifera, Central Apennines.

Riassunto - I depositi pre-terrigeni del Miocene superiore dell'unità Queglia (Abruzzo, Appennino centrale, Italia): dati biostratigrafici. - È stato condotto uno studio biostratigrafico integrato (nannofossili calcarei, foraminiferi planctonici e bentonici) sui depositi pre-terrigeni alto miocenici dell'unità Queglia. La sezione studiata, situata nei pressi dell'abitato di Pescosansonesco (Abruzzo, Italia), è costituita da depositi di rampa deepening-upward d'avampaese, che durante il Messiniano occupava una posizione più esterna rispetto al ben noto bacino d'avanso della Laga.

La presenza di *Amaurolithus primus*, e di *Globorotalia, gr. miotumida* è stata registrata sin dalla base della sezione, dove si assiste ad un passaggio graduale da sedimenti di mare basso a sedimenti più distali. Nei sovrastanti depositi emipelagici sono state registrate la FO di *Bulimina echinata*, le FO e FCO di *Turborotalia multiloba*, e la FO di *Amaurolithus amplificus*.

Questi bioeventi hanno consentito di riconoscere la Subzona *Turborotalia multiloba* e la Subzona *Calcidiscus leptoporus* MA.

Parole chiave - Biostratigrafia, Nannofossili calcarei, Foraminiferi, Appennino centrale.

INTRODUCTION

The Queglia unit, outcropping North and West of the Maiella mountain, belongs to a paleogeographic

domain located between the Laga (inner) and Maiella (outer) domains (Fig. 1).

According to Patacca *et al.* (1991) the Queglia unit consists of Cretaceous-Paleogenetic basinal carbonates, evolving to outer ramp deposits in the late Oligocene? - Middle Miocene and to inner ramp deposits in the early Tortonian. During the early Messinian the deposition of hemipelagic sediments testified to the drowning of the carbonate platform. As in many Mediterranean sequences the «Lower evaporitic Cycle» was registered and above this the «Lago Mare» deposits were recorded. In the latest Messinian the onset of the synorogenic turbidite siliciclastic deposition («flysch Teramano») occurred.

The object of this study is the biostratigraphy of the hemipelagic deposits marking the change from carbonate to terrigenous sedimentation.

Patacca *et al.* (1991) assign these sediments to the early Messinian, on the basis of the occurrences of *Amaurolithus primus*, *Amaurolithus delicatus*, *Amaurolithus amplificus*, *Turborotalia multiloba* and *Bulimina echinata*.

The aim of this study is to provide further information about the calcareous benthic and planktic assemblages of these sediments.

LITHOSTRATIGRAPHY

The section studied is located near the Pescosansonesco village (Lat. 46°77'27.031; Long. 40°74'43.958; F°146-II NE) along the road to the village of Corvara (Fig. 2).

The interval sampled is about 40 meters thick starting from the boundary with the underlying carbonate platform sediments of the Bolognano formation. Four different superposed lithofacies are recognized (Fig. 3).

The lower part of the section (lithofacies 1) is made up of about 6 metres of grey glauconitic, bioturbated marly-limestones with abundant macrofossils, and thin interlayers of brownish-grey marls with bivalves and planktic foraminifera.

The overlaying hemipelagic sediments consist of 4 metres of grey marls and clayey marls with pteropods (lithofacies 2).

Upsection, the marls with pteropods give way to a marly

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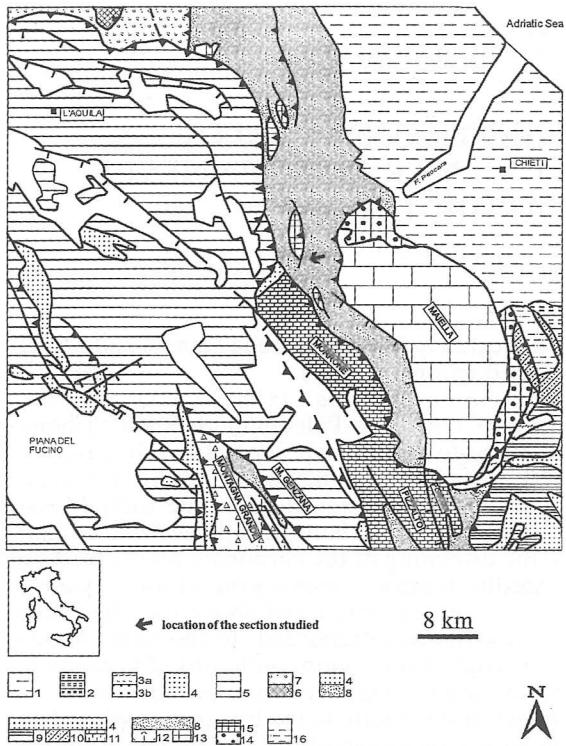


Fig. 1 - Geologic-structural scheme of the Central Apennines-Abruzzo (Patacca *et al.* 1991, modified).

1) Upper Pliocene p.p.-Pleistocene cycle; 2) Upper Pliocene p.p. cycle; 3a) Messinian-«Lagomare»-Lower Pliocene cycle; 3b) Palena and La Vicenne conglomerates; 4) foredeep deposits of the pre-evaporitic early Messinian: «flysch abruzzesi» and «flysch molisani» (Cantalupo, S.Massimo, S.Elena, and Agnone flysch, Olmi formation and Treste formation); 5) pre-terrigenous deposits of Gran Sasso and eastern Marsica units; 6) pre-terrigenous deposits of Montagna dei Fiori unit; 7) foreland deposits of pre and post evaporitic Messinian (Laga flysch); 8) foredeep deposits of the Messinian «Lagomare»-Early Pliocene p.p.; 9) Agnone; 10) Tuffillo; 11) pre-terrigenous deposits of Scontrone-Porrara unit; 12) pre-terrigenous deposits of the Montagna Grande; 13) pre-terrigenous deposits of the Queglia unit; 14) pre-terrigenous deposits of the Maiella unit; 15) Lower Pliocene p.p. foredeep deposits; 16) foreland pelitic deposits of the Casoli-Bomba unit (Late Pliocene p.p.-Early Pliocene).

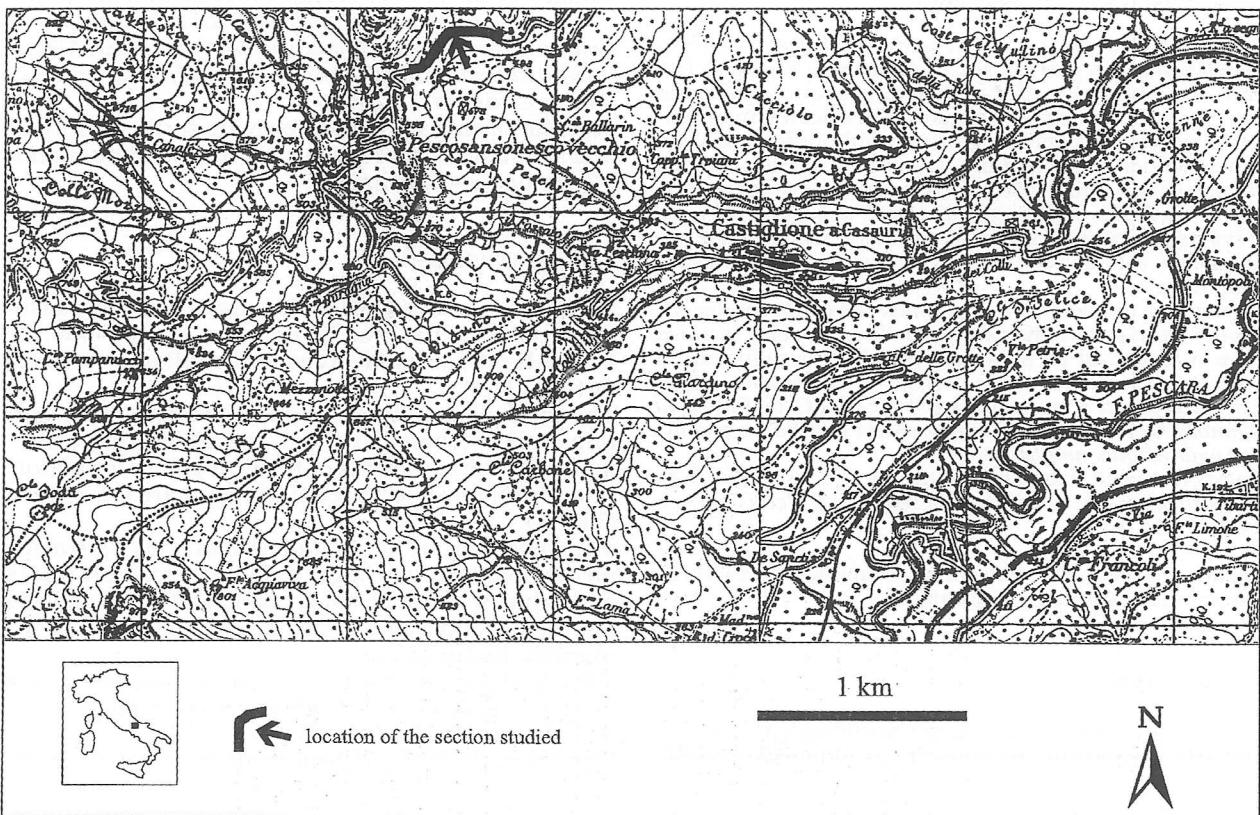


Fig. 2 - Location map of the Pescosansonesco section (from the IGM Carta d'Italia, scale 1:25.000, sheet 146 I SE, Popoli).

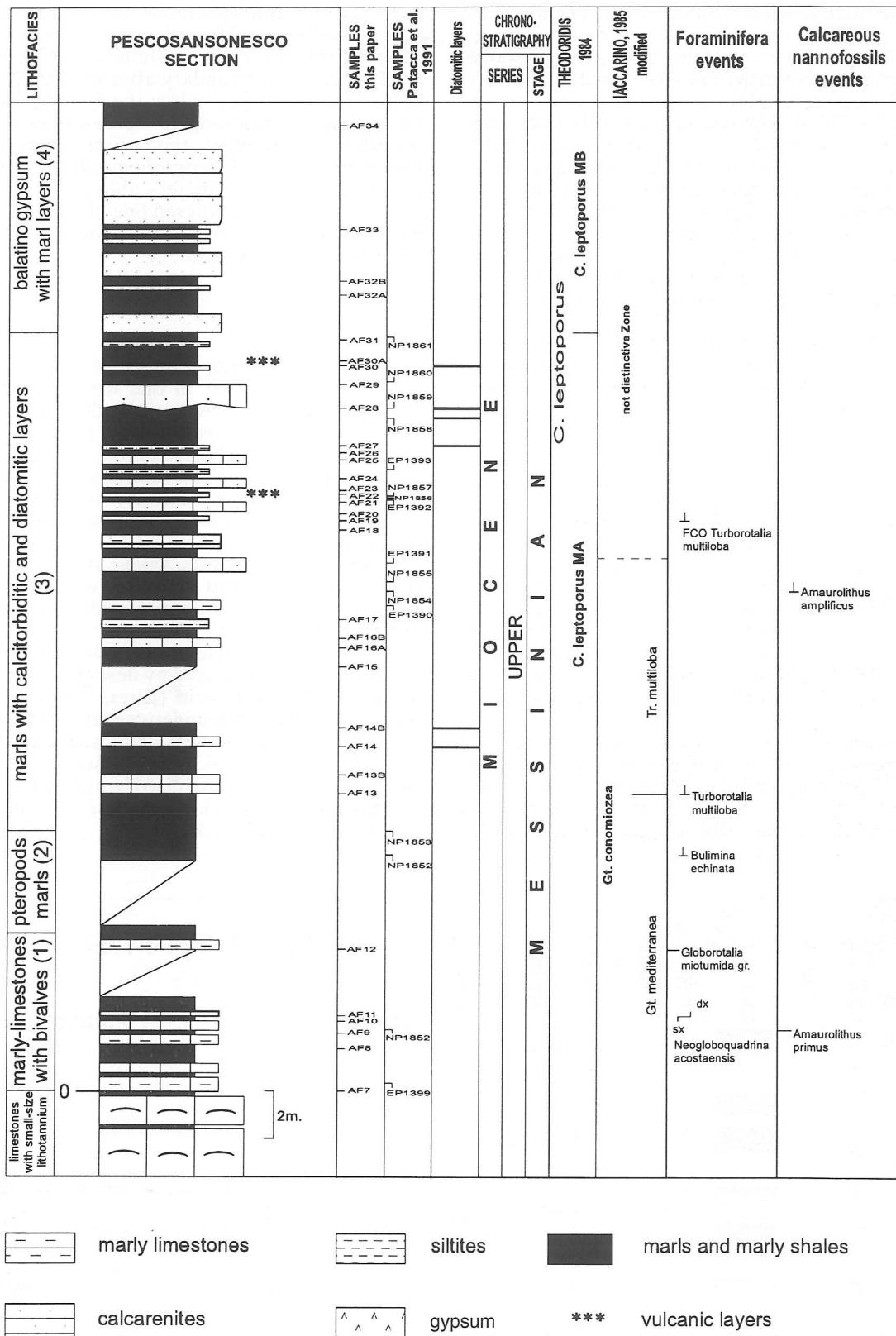


Fig. 3 - Lithostratigraphy, chronostratigraphy, biochronology and bioevents of the Pescosansonesco section.

18 metre thick interval (lithofacies 3). The lower and middle portions are characterized by an irregular alternation of greenish-grey marls rich in sponge spicules, greenish-grey lime-marls with planktic foraminifera, rare thin brownish siltites, and very rare dark shales with bivalves (sample AF16B). In the upper part of this lithofacies an increase of calciturbiditic layers is observed, characterized by medium to coarse-grained bioclastic calcarenites interlayered with greenish-grey marls, lime-marls with planktic foraminifera, and brownish siltites.

In the upper part of lithofacies 3 two volcaniclastic levels were recognized. Some diatomitic layers were recorded both in the lower and in the upper part of this lithofacies.

The section ends with about 10 metres of balatino gypsum (lithofacies 4). In the lower portion of this interval intercalations of green marls and rare grey siltites occur.

BIOSTRATIGRAPHY

The biostratigraphic scheme used in this study for the Messinian interval is shown in Fig. 4.

For the planktic foraminifera biostratigraphy, the Mediterranean scheme of Iaccarino 1985 (modified according to Colalongo *et al.*, 1979, and Sprovieri *et al.*, 1996a and 1996b) is applied. The calcareous nannofossil zonation is based on the Mediterranean scheme of Theodoridis (1984) correlated to the zonal scheme of Martini (1971) and Okada & Bukry (1980).

The calcareous plankton events are recalibrated on the astronomical and geomagnetic time scale of Cande & Kent (1995) according to Sprovieri *et al.* (1996b). Remarks - The base of «non distinctive Zone» (Iaccarino, 1985) is normally recognized when essen-

tially right coiling specimens of *Neogloboquadrina acostaensis* occur. However, due to the scarce presence of this taxon in the studied section we tentatively locate this boundary after the FO of *Amaurolithus amplificus* just below the FCO of *Turborotalia multiloba* according to Sprovieri *et al.* (1996a and 1996b). Therefore, the coiling change from sinistral to dextral of *Neogloboquadrina acostaensis* recorded at about 4 meters above the base of the section could be considered one of the coiling changes of this species during early Messinian (Iaccarino, 1998, pers. comm.).

METHODS

49 samples were analyzed to study foraminifera assemblage (33 samples in this work, 16 samples in Patacca *et al.*, 1991). As regards the calcareous nannofossils 35 samples were studied (33 samples in this work, 2 in samples Patacca *et al.*, 1991).

In each sample, the foraminifera and nannofossil assemblages were qualitatively and semiquantitatively characterized in terms of preservation and abundance. (Tab. 1)

For nannofossil analysis, smear slides of each sample were prepared and examined with a light microscope at $\times 1200$ magnification under cross-polarized and transmitted light.

The foraminifera assemblages were studied from washed residues produced by desegregating the rock-samples with acetic acid (Lirer, 1998) and sieving the residues. Residues underwent ultrasonic treatment in order to obtain cleaner residues. The analysis was carried out on the $> 63 \mu\text{m}$ fraction.

In Tables 2, 3 and 4 the distribution of foraminifera (planktic and benthic), calcareous nannofossils and ostracods were plotted.

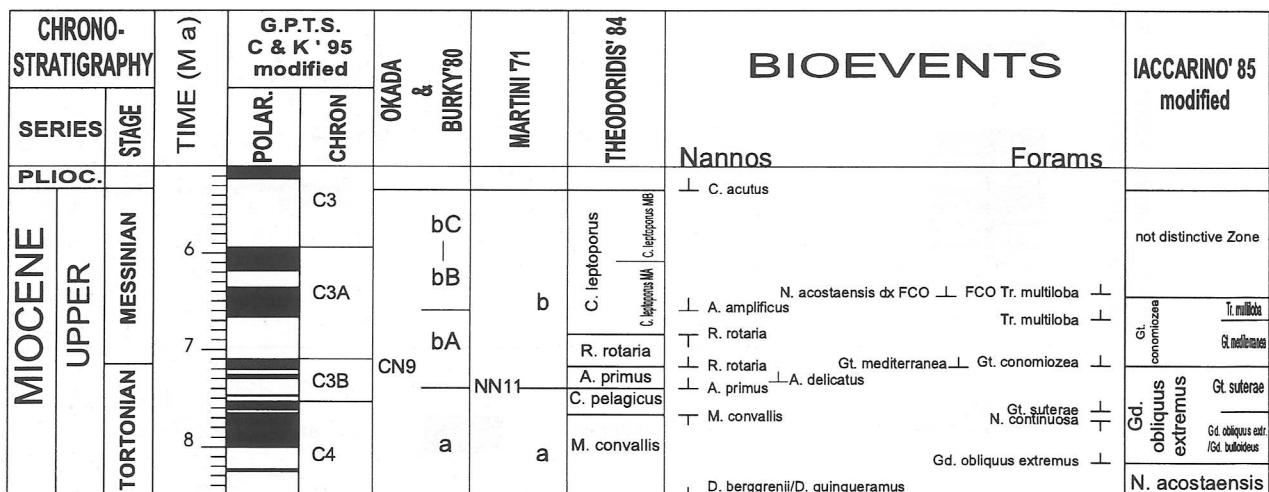


Fig. 4 - Adopted Late Miocene integrated biostratigraphic scheme according to: Cande & Kent (1995), Sprovieri *et al.* (1996b), (time scale and magnetic calibration); Shackleton *et al.* (1995) Sprovieri *et al.* (1996b), (calcareous nannofossils). Sprovieri *et al.* (1996b), (planktic foraminifera).

Tab.1

| CALCAREOUS NANNOFOSSILS | FORAMINIFERA |
|--|---|
| <p>TOT. AB. A = > 30 spec.s for view; C = 10 to 30 spec.s for view; F = 1 to 10 spec.s for view; R = < 0.1-1 spec.s for view; P = < 0.1 spec.s for view; B = barren</p> <p>REL. AB. A = > 1 spec for view; C = 1 spec for view; F = 1 spec for 2-10 views; R = 1 spec for 10-100 views; P = 1 spec for 100-600 views</p> <p>PRESERVATION: G (good): little or no evidence of dissolution and/or secondary overgrowth; fully preserved diagnostic characters. M (moderate): dissolution and/or secondary overgrowth; partially altered primary morphological characteristics; nearly all specimens can be identified at the species level. P (poor): severe dissolution, fragmentation and/or secondary overgrowth; largely destroyed primary features; many specimens cannot be identified at the species and/or at generic level.</p> | <p>TOT. AB. A > 30%; C = 10-30%; F = 3-10%; R = 1-3%; P = less than 1%; B = barren</p> <p>REL. AB. A > 30%; C = 10-30%; F = 3-10%; R = 1-3%; P = less than 1%</p> <p>PRESERVATION: G (good): little or no fragmentation, overgrowth and/or dissolution (> 90% of intact specimens); M (moderate): some signs of fragmentation, overgrowth and/or dissolution (30 to 90% of intact specimens); P (poor): severe fragmentation, heavy overgrowth and/or dissolution (< 30% of intact specimens)</p> |

RESULTS

Foraminifera

The samples studied yielded a moderately preserved, low to high diversified and abundant foraminifera assemblages. Only a few samples, collected in the diatomitic layers, contained a very poor foraminifera fauna.

In the lowermost 11 metres of the section, the planktic assemblage is rich and moderately preserved, being characterized by: *Globorotalia miotumida* gr., *Globigerinella obesa*, *Neogloboquadrina acostaensis* sx., *N. acostaensis* dx., *Globigerinoides obliquus extremus*, *Gd. obliquus obliquus*, *Globorotalia humerosa praehumerosa*, *Hastigerina siphoniphera*, *Zeaglobigerina apertura*, *Zg. woodi woodi*, *Zg. nepenthes*, *Zg. decoraperta*, *Globigerina bulloides*, *Orbulina suturalis* and *O. universa* (Table 2).

With regard to the benthic foraminifera, the association consists of *Brizalina* sp., *B. aff. dilatata*, *B. dentellata*, *Rectuvigerina siphogenerynoides*, *R. gaudryinoides*, *Heterolepa* sp., *H. mexicana*, *Elphidium crispum*, *Uvigerina* sp., *Cancris oblungus*, *Gyroidina* sp., *Gyroidinoides* sp., *G. altiformis*, and rare *Cibicidoides pachyderma* (Table 3). The first occurrence of *Bulimina echinata* (sample NP 1852, Tab. 2 and Fig. 5) was recorded at about 10 metres from the base of the section.

In this portion of the section, the coiling change of *Neogloboquadrina acostaensis* (sample AF11, Tab. 3 and Fig. 3) and the appearance of *Globorotalia gr. miotumida* (sample AF12, Tab. 3) were also recorded. We did not consider the appearance of *Globorotalia gr. miotumida* as its FO, because it occurred when the area studied underwent a change from a shallow to deeper water shelf. To support our hypothesis we propose a stratigraphic correlation with the Pietrasecca section (Cosentino *et al.*, 1997, Pampaloni *et al.*, 1994), occupying an inner domain with

respect to the Queglia unit. In the Pietrasecca section the FO of *Globorotalia gr. miotumida* occurs in the upper part of the «Marne ad Orbulina» formation, which represents the involvement of Pietrasecca area in a foredeep system (foreland flexure stage) while in the Queglia domain the sedimentary record was still represented by foreland shallow water carbonate deposits (Scandone, 1998, pers. comm.). Throughout the remaining portion of the section the organic content is poorly diversified both in benthic and planktic foraminifera (Fig. 5).

In sample AF13 the FO of *Turborotalia multiloba* (Tab. 3 and Fig. 5) was recorded; this allowed the recognition of the *Turborotalia multiloba* Subzone (Iaccarino, 1985, modified). After this biostratigraphic event the planktic and benthic foraminifera assemblages become progressively oligotypical.

The planktic foraminifera assemblage is characterized by *Orbulina suturalis*, *O. universa*, *Turborotalia quinqueloba*, *Zeaglobigerina decoraperta*, *Zg. apertura*, *Globigerina bulloides* and rare *Neogloboquadrina acostaensis* sx.

The benthic foraminifera assemblage is characterized by *Bulimina aculeata minima*, *B. echinata*, *Brizalina* sp., *B. aff. dilatata*, *B. dentellata*, *Elphidium crispum* and *Cibicidoides* sp.. *Rectuvigerina siphogenerynoides*, *R. gaudryinoides* are present only in the lowermost part of this interval. This benthic assemblage corresponds to the one recorded in the Tripoli Fm by Colalongo *et al.* (1979).

11 metres below the evaporitic episode, the FCO of *Turborotalia multiloba* was registered; this is in substantial agreement with the data of Sprovieri *et al.* (1996a and 1996b).

Some horizons (Fig. 5 and Tab. 2) with planktic association characterized by the 80-90% of *Orbulina*, were recognized. These levels are very similar to the «Orbuliniti», described in the upper part of the Tripoli Formation (Sprovieri *et al.*, 1996a and 1996b, Colalongo *et al.*, 1979).

Table 2 - Distribution of planktic foraminifera in the Pescosansonesco section.

The organic content decreases near the evaporitic episode (lithofacies 4, Fig. 5), and the marls interlayered with the balatino gypsum are completely barren (Fig. 5). Marine ostracods of inner neritic environment are present in some horizons of the pteropods marls in-

terval and marls with diatomitic layers in the specimens: *Cytherelloidea* cf. *cretensis*, *Loxoconcha punctatella*, *Xestoleberis* sp., *Aurila philippii*, *A. convexa*, *A. cf. impressa*, *A. freudentali*. (specific identification by B. Dell'Antonia; Tab. 3)

Table 3 - Distribution of benthic foraminifera in the Pescosansonesco section.

| LITHOFACIES | | SAMPLES | | Total abundance | Indeterminate planktic forams | | | | | | | | | | | | | | | PLANK. FORAM. ZONE (Iaccarino, 1985 modified) | |
|---|--|---------|-------------|-----------------|----------------------------------|--|--|--|--|--|--|--|--|--|--|--|--|--|--|--|--|
| | | | | | Preservation | | | | | | | | | | | | | | | | |
| balatino gypsum with marl layers (4) | | AF 34 | B | | <i>Globigerinoides</i> sp. | | | | | | | | | | | | | | | | |
| | | AF 33 | B | | <i>G. praebulloides</i> | | | | | | | | | | | | | | | | |
| | | AF 32B | B | | <i>D. globularis</i> | | | | | | | | | | | | | | | | |
| | | AF 32A | B | | <i>Globigerinoides</i> sp. | | | | | | | | | | | | | | | | |
| | | NP 1861 | F P R | | <i>P. siakensis</i> | | | | | | | | | | | | | | | | |
| | | AF 31 | B | | <i>Ge. obesa</i> | | | | | | | | | | | | | | | | |
| marls with calciturbiditic and diatomitic layers (3) | | AF 30A | F P | | <i>Gq. baromoenensis</i> | | | | | | | | | | | | | | | | |
| | | AF 30 | B | | <i>Gd. quadrilobatus</i> | | | | | | | | | | | | | | | | |
| | | NP 1860 | A M | | <i>Zg. woodi woodi</i> | | | | | | | | | | | | | | | | |
| | | AF 29 | B | | <i>G. falconensis</i> | | | | | | | | | | | | | | | | |
| | | AF 28 | B | | <i>Gd. trilobus</i> | | | | | | | | | | | | | | | | |
| | | NP 1859 | F M-P | | <i>D. trilobata</i> | | | | | | | | | | | | | | | | |
| | | NP 1858 | C M F P | P | <i>Gt. acrostoma</i> | | | | | | | | | | | | | | | | |
| | | AF 27 | B | | <i>Ge. praesiphoniphera</i> | | | | | | | | | | | | | | | | |
| | | AF 26 | C P F F | | <i>P. glomerosa</i> | | | | | | | | | | | | | | | | |
| | | EP 1393 | | | <i>O. sibularis</i> | | | | | | | | | | | | | | | | |
| | | AF 24 | F P | | <i>O. bilobata</i> | | | | | | | | | | | | | | | | |
| | | AF 23 | B | | <i>O. universa</i> | | | | | | | | | | | | | | | | |
| | | AF 22 | F P F | | <i>Gd. bollii</i> | | | | | | | | | | | | | | | | |
| | | NP 1857 | A M F | | <i>Gd. obliquus obliquus</i> | | | | | | | | | | | | | | | | |
| | | NP 1856 | A M | | <i>Zg. decorperita</i> | | | | | | | | | | | | | | | | |
| | | EP 1392 | C P F R | | <i>Tr. quinqueloba</i> | | | | | | | | | | | | | | | | |
| | | AF 21 | C M F | | <i>Zg. nepenthes</i> | | | | | | | | | | | | | | | | |
| | | AF 20 | C M F | | <i>Gt. gr. menardii</i> | | | | | | | | | | | | | | | | |
| | | AF 19 | C M-P F F | | <i>Gt. partimlabiata</i> | | | | | | | | | | | | | | | | |
| | | AF 18 | F P | | <i>Gt. merotumida</i> | | | | | | | | | | | | | | | | |
| | | NP 1855 | B | | <i>Zg. apertura</i> | | | | | | | | | | | | | | | | |
| | | NP 1854 | C M-P R F | R | <i>G. bulloides</i> | | | | | | | | | | | | | | | | |
| | | EP 1390 | R P R | | <i>N. acostensis</i> ss | | | | | | | | | | | | | | | | |
| | | AF 17 | B | | <i>Ge. siphoniphera</i> | | | | | | | | | | | | | | | | |
| | | AF 16B | R P R | | <i>Gd. obliquus extremus</i> | | | | | | | | | | | | | | | | |
| | | AF 16 | F M-P F | | <i>Gt. humerosa praehumerosa</i> | | | | | | | | | | | | | | | | |
| | | AF 15 | C M-P F | | <i>Gt. miotumida</i> | | | | | | | | | | | | | | | | |
| | | AF 14B | B | | <i>Zg. apertura</i> | | | | | | | | | | | | | | | | |
| | | AF 14 | P P R P | | <i>G. bulloides</i> | | | | | | | | | | | | | | | | |
| | | AF 13B | A M-P F R | R R | <i>N. acostensis</i> ss | | | | | | | | | | | | | | | | |
| | | AF 13 | C P F F | | <i>Ge. siphoniphera</i> | | | | | | | | | | | | | | | | |
| pteropods marls (2) | | NP 1853 | A M-P C | P F F | R C A | | | | | | | | | | | | | | | | |
| | | NP 1852 | C M-P F R | P R P | R R P | | | | | | | | | | | | | | | | |
| | | AF 12 | C P F F | F | F F C F | | | | | | | | | | | | | | | | |
| | | AF 11 | C P F | R | F | | | | | | | | | | | | | | | | |
| | | AF 10 | A M-P C | | | | | | | | | | | | | | | | | | |
| | | NP 1851 | C P F R | | | | | | | | | | | | | | | | | | |
| | | AF 9 | A M F C | R P | C R F F | | | | | | | | | | | | | | | | |
| | | AF 8 | A M-P F C R | F R | R F C C F | | | | | | | | | | | | | | | | |
| | | EP 1389 | C P C F | R | F C F P R | | | | | | | | | | | | | | | | |
| | | AF 7 | A M-P C | R | C C R | | | | | | | | | | | | | | | | |
| | | | | | | | | | | | | | | | | | | | | | |
| marly limestones with bivalves (1) | | | | | | | | | | | | | | | | | | | | | |
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Nannofossils

The nannofossils are generally abundant and moderately preserved throughout the section studied (Tab. 4). The assemblage consists mainly of abundant to common small-sized morphotypes of genera *Reticu-*

lofenestra and *Dictyococcites*. *Calcidiscus leptopus*, *Helicosphaera carteri*, *Sphenolithus* spp., *Syracosphaera* spp., *Scyphosphaera* spp. are commonly recorded as well. *Discoaster* spp. (five and six rayed morphotypes), *Amaurolithus primus*, *A. delicatus*, *A. amplificus* (whose FO was detected at about 27 meters

Turborotalia multiliba subzone
Globorotalia mediterranea subzone

non distinctive Zone

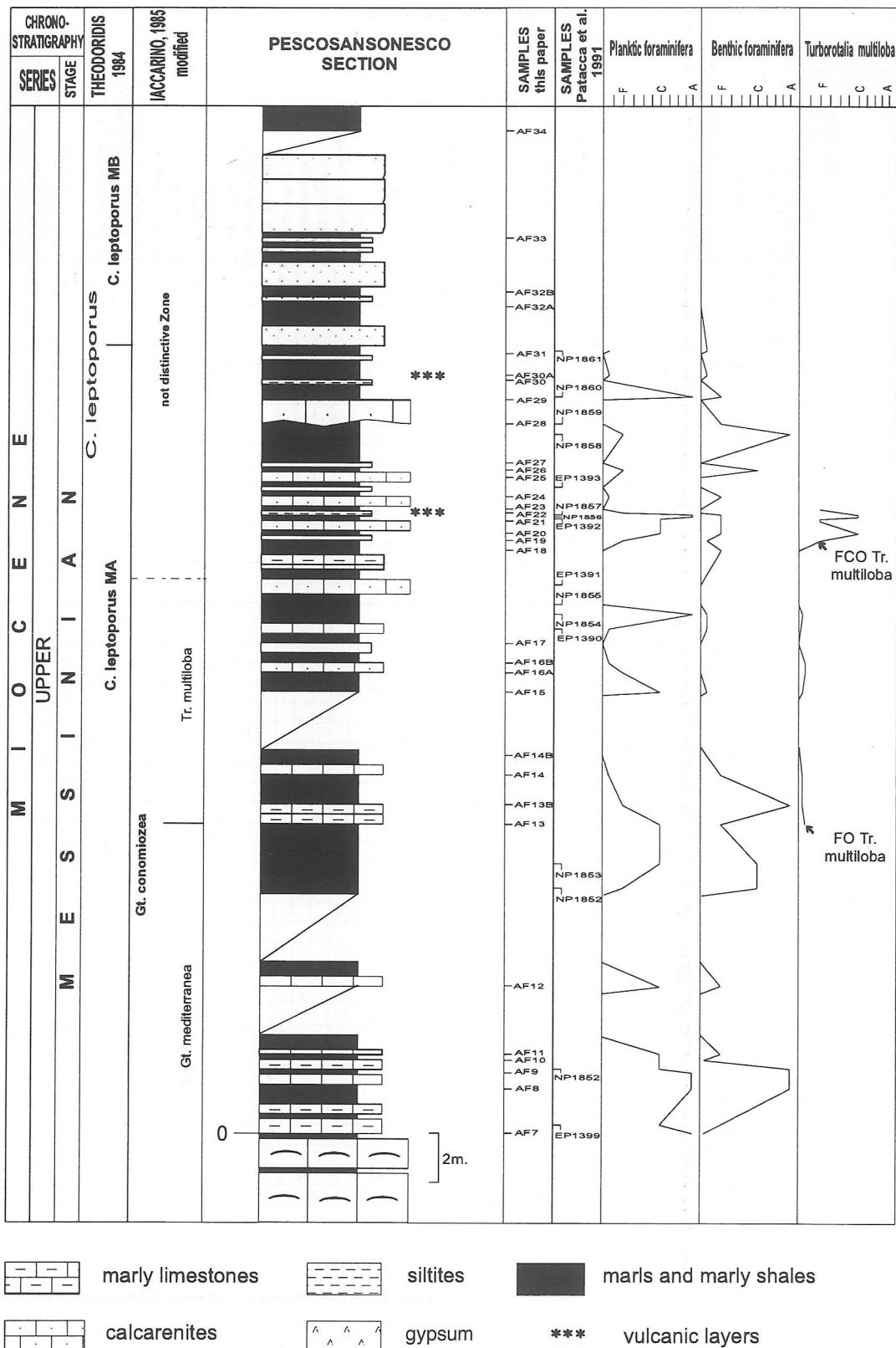


Fig. 5 - Abundance patterns of planktic, benthic foraminifera and *Turborotalia multiloba* in the Pescosansonesco section.

Table 4 - Distribution of calcareous nannofossils in the Pescosansonesco section.

from the base of the section, sample NP1854), *Calcidiscus macintyreai*, *Geminilitella rotula* and *Braarudosphaera bigelowii* occur in lower percentage. Common Late Cretaceous, Paleogene and Early Miocene reworked specimens are also present. For this reason, we place the section studied within the *Calcidiscus leptoporus* MA Subzone (Theodoridis, 1984), due to the occurrences of *Amaurolithus primus*, *Amaurolithus delicatus*, *Amaurolithus amplificus* and to the absence of *Reticulofenestra rotaria*. The abundance of Discoasterids is generally low throughout the section; noteworthy are the levels rich in siliceous microfossils and organic matter, that are devoid of Discoasterids. According to Raffi & Flores (1995) the occurrence of Discoasterid-depleted levels may be linked to enhanced upwelling conditions. Moreover, the state of preservation of *Discoaster* is so poor that specific identification is hampered; for this reason we prefer to label them only as *Discoaster* spp. six and/or five rayed forms.

The genus *Helicosphaera* is represented mainly by *Helicosphaera carteri*, occurring with high-frequency values (almost 90%); we were not able to detect the occurrence of «small» *Helicosphaera* whose LO follows the FO of *Amaurolithus* (Negri & Vigliotti, 1997).

The occurrence of *Amaurolithus* is scattered and with low-frequency values. We recorded the presence of *Amaurolithus primus*, *Amaurolithus delicatus* and the FO of *Amaurolithus amplificus*. The presence of many intermediate forms is noteworthy suggesting the possibility of more detailed analysis of this genus in this interval.

As for the occurrence of *Globorotalia miotumida* gr. (see previous paragraph), we do not consider the appearance of *Amaurolithus primus* in the lower part of the section as its first occurrence.

In the marls with diatomitic layers an oligotypical assemblage dominated by *Calcidiscus leptoporus* was recorded (sample AF 13). Similar levels were described from the Tripoli Formation by Sprovieri *et al.*, (1996a and 1996b).

CONCLUSIONS

This study provides new information about the early Messinian bioevents (calcareous nannofossils and foraminifera) in the hemipelagic deposits of the Queglia unit.

We were able to state more exactly the FO of *Bulimina echinata*, FO and FCO of *Turborotalia multiloba* and FO of *Amaurolithus amplificus*.

Noteworthy are the many analogies between the Pescosansonesco section and the typical lower Messinian Mediterranean sequences (Sprovieri *et al.*, 1996a and 1996b; Colalongo *et al.*, 1979) such as:

- the appearance of *Bulimina echinata* marking a change from highly diversified assemblages to oligotypical ones both for planktic and benthic foraminifera;
- the presence of «Orbuliniti» and diatomitic layers;

– the presence of oligotypical calcareous nannofossils assemblages dominated by *Calcidiscus leptoporus*;

– the FO of *Amaurolithus amplificus* predating the FCO of *Turborotalia multiloba* just below the evaporitic episode.

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TAXONOMIC REFERENCES

Selected foraminifera and calcareous nannofossils considered in this study (in alphabetic order of generic epithets).

Planktic Foraminifera

Globigerina bulloides d'Orbigny, 1826

Globigerinoides obliquus extremus Bolli & Bermudez, 1965

Globigerinoides obliquus obliquus Bolli, 1957

Globorotalia humerosa praehumerosa Natori, 1976

Globorotalia gr. *miotumida* sensu Sierro *et al.*, 1993

PLATE 1

- | | |
|--------------|---|
| Figs. 1, 2 | <i>Globigerinoides obliquus extremus</i> Bolli & Bermudez. Fig. 1: umbilical view; Fig. 2: spiral view. Sample AF8. (x 60). |
| Figs. 3, 4 | <i>Globigerinoides obliquus obliquus</i> Bolli. Fig. 3: umbilical view; Fig. 4: spiral view. Sample AF8. (x 60). |
| Figs. 5, 6 | <i>Globorotalia miotumida</i> gr. (sensu Sierro <i>et al.</i> , 1993). Fig. 5: umbilical view; Fig. 6: axial view. Sample AF12. (x 60). |
| Figs. 7, 8 | <i>Globorotalia humerosa praehumerosa</i> Natori. Fig. 7: umbilical view; Fig. 8 spiral view. Sample AF9. (x 60). |
| Figs. 9, 10 | <i>Neoglobiquadrina acostaensis</i> (Blow) dx. Fig. 9: umbilical view; Fig. 10: spiral view. Sample AF11. (x 110). |
| Figs. 11, 12 | <i>Neoglobiquadrina acostaensis</i> (Blow) sx. Fig. 11: umbilical view; Fig. 12: spiral view. Sample AF11. (x 110). |
| Figs. 13, 14 | <i>Turborotalia multiloba</i> (Romeo). Fig. 13: umbilical view; Fig. 14: spiral view. Sample AF13. (x 140). |
| Figs. 15, 16 | <i>Turborotalia quinqueloba</i> (Natland). Fig. 15: umbilical view; Fig. 16: spiral view. Sample AF19. (x 140). |
| Fig. 17 | <i>Brizalina arta</i> (Mac Fadyen). Sample AF9. (x 40). |
| Fig. 18 | <i>Bulimina aculeata minima</i> Tedeschi & Zanmatti. Sample AF13. (x 110). |
| Fig. 19 | <i>Bulimina echinata</i> d'Orbigny. Sample AF13. (x 110). |
| Fig. 20 | <i>Rectuvigerina gaudryinoides</i> (Lipparini). Sample AF8. (x 25). |
| Fig. 21 | <i>Rectuvigerina siphogenerinoides</i> (Lipparini). Sample AF8. (x 30). |
| Fig. 22 | <i>Brizalina dentellata</i> (Tavani). Sample AF20. (x 60). |

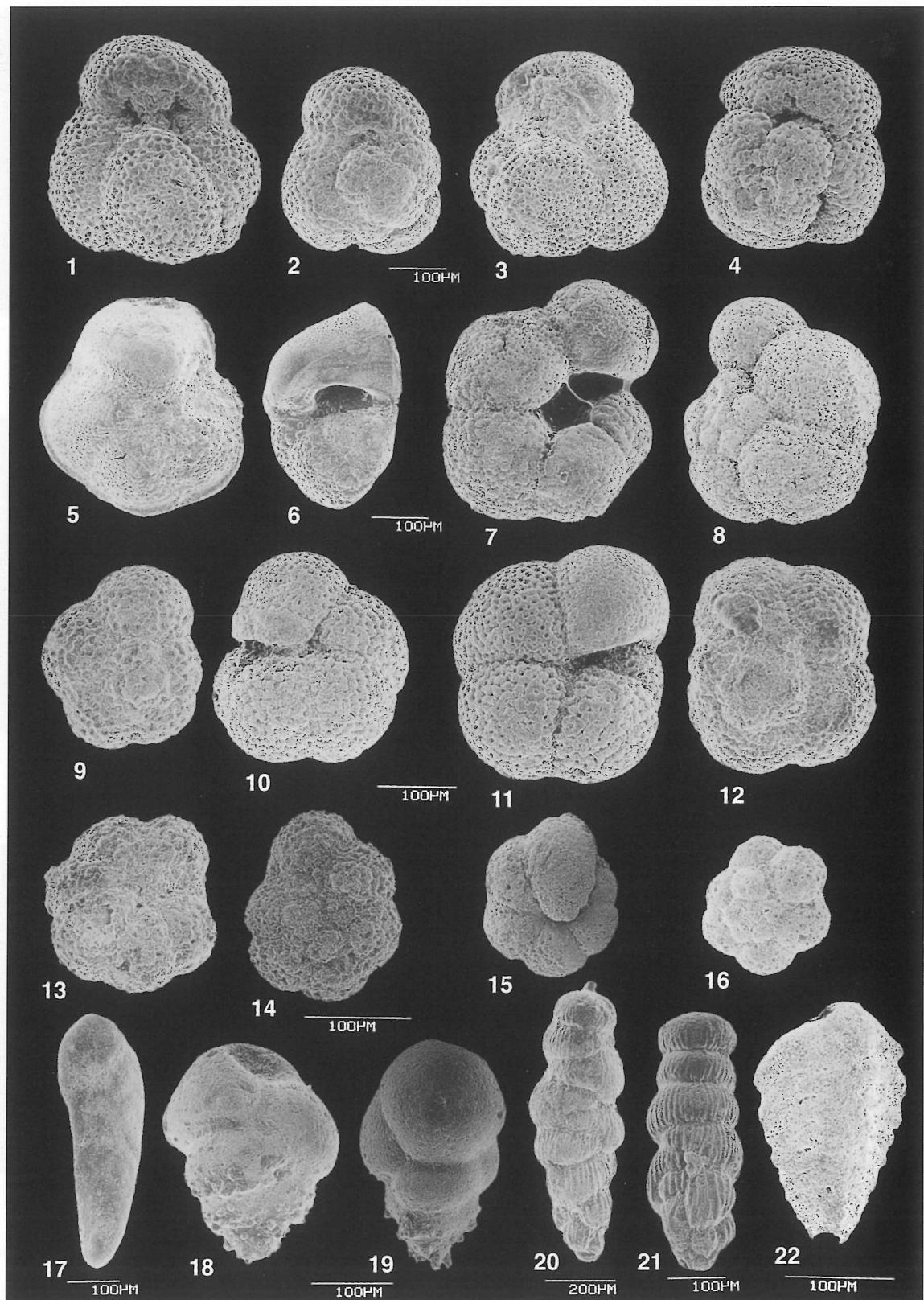
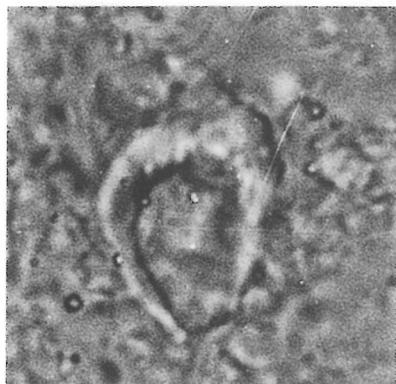


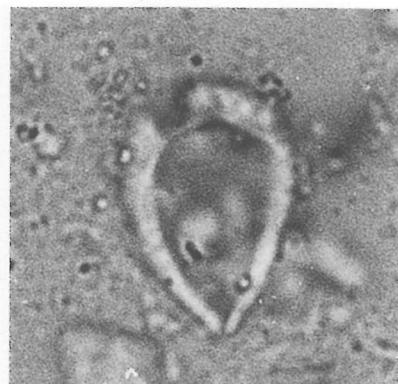
PLATE 1



1



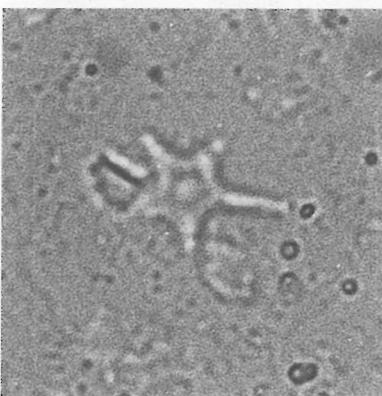
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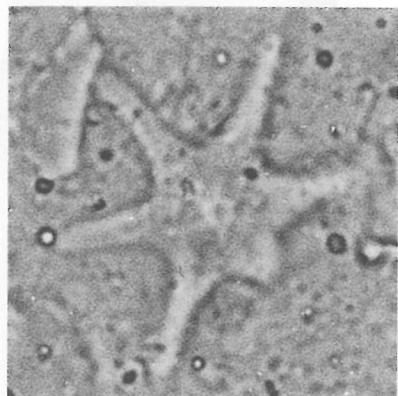
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4



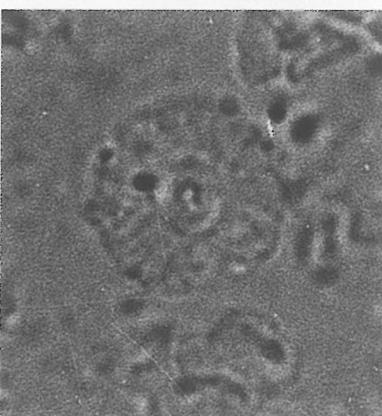
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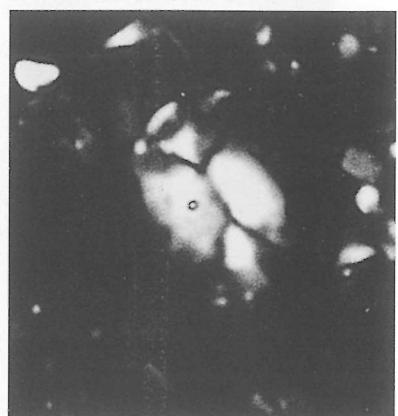
6



7



8



9

PLATE 2

All specimens x 2400. Fig. 1: *Amaurolithus primus* (Bukry and Percival) Gartner and Bukry. Sample AF9A. Fig. 2: *Amaurolithus* sp.. Sample AF13B. Fig. 3: *Amaurolithus delicatus* Gartner & Bukry. Sample AF23. Fig. 4: *Amaurolithus* sp.. Sample AF23. Fig. 5: five-rayed *Discoaster* sp., sample AF13. Fig. 6: six-rayed *Discoaster* sp.. Sample AF13. Fig. 7: *Amaurolithus amplificus* (Bukry & Percival) Gartner and Bukry. Sample NP1854. Fig. 8: *Calcidiscus leptoporus* (Murray & Blackman) Loeblich & Tappan. Sample AF13B. Fig. 9: *Helicosphaera carteri* (Wallich) Kamptner. Sample AF 13.

Hastigerina siphoniphera siphoniphera (d'Orbigny), 1839
Neogloboquadrina acostaensis acostaensis (Blow), 1959
Orbulina suturalis Brönnimann, 1951
Orbulina universa d'Orbigny, 1839
Turborotalia multiloba (Romeo), 1965
Turborotalia quinqueloba (Natland), 1938
Zeaglobigerina apertura (Cushman), 1918
Zeaglobigerina decoraperta (Takayanagi & Saito), 1962
Zeaglobigerina nepenthes (Tood), 1957
Zeaglobigerina woodi woodi (Jenkins), 1960
Benthic Foraminifera
Brizalina arta (MacFadyen), 1930
Brizalina dentellata (Tavani), 1955
Brizalina aff. Brizalina dilatata (Reuss, 1850)
Bulimina aculeata minima Tedeschi & Zanmatti, 1957
Bulimina echinata d'Orbigny, 1852
Cancris oblongus (Williamson), 1858
Cibicidoides pachyderma (Rzehak), 1886
Elphidium crispum (Linné), 1758
Gyroidinoides altiformis (Stewart and Stewart), 1930
Heterolepa mexicana (Nuttall), 1932
Rectuvigerina gaudryinoides (Lipparini), 1932
Rectuvigerina siphogenerinoides (Lipparini), 1932
Calcareous Nannofossils
Amaurolithus amplificus (Bukry and Percival, 1971)
Gartner and Bukry, 1975
Amaurolithus delicatus Gartner and Bukry, 1975
Amaurolithus primus (Bukry and Percival, 1971)
Gartner and Bukry, 1975
Braarudosphaera bigelowii (Gran and Braarud, 1935)
Deflandre, 1974
Calcidiscus leptoporus (Murray and Blackman, 1988)
Loeblich and Tappan, 1978
Calcidiscus macintyreai (Bukry and Bramlette, 1969)
Loeblich and Tappan, 1978
Geminolithella rotula (Kamptner, 1956). Backman, 1980
Helicosphaera carteri (Wallich, 1877) Kamptner, 1954
Reticulofenestra rotaria Theodoridis, 1984

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